

Active source seismic studies across/along across-arc seamount chains in Izu-Bonin arc - Cruise report of KR0601 and KR0605 -

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Abstract In order to investigate formation processes of the across-arc sea mount chains in the Izu-Bonin arc, we have carried out two active source seismic cruises by using JAMSTEC R/V Kairei in 2006. The wide-angle seismic data were acquired across the across-arc seamount chains (Nishi-shichito ridge) in KR0605 and along the across-arc seamount chain (Minami-Iojima ridge) in KR0601. In addition to those two wide-angle profiles, multichannel seismic data were also acquired along two profiles which are parallel to the strike of the across arc seamount chains. The wide-angle seismic data indicate thinning of crust toward the south across the Nishi-shichito ridge, and relatively flat and thin crust beneath the Minami-Ioujima ridge. The multi-channel data show that the basement of the Shikoku basin can be traced beneath the western flank of the across-arc seamount chain. This suggest that a boundary between the Shikoku basin crust and the Oligocene Izu arc may be situated around the eastern end of the across-arc seamount chains.

Keywords: Crustal structure, active source seismic study, Izu-Bonin, across-arc seamount chains

1. Introduction

The Izu-Bonin-Mariana (IBM) intra-oceanic arc system extends over 2800 km from Sagami Bay in the north to Guam in the south, and has long been studied as an excellent example of an intra-oceanic island arc system. The tectonic history of the IBM arc is well documented in many studies. Here, we briefly summarize those studies of relevance to ours, following the work of Stern et al. (2003). Estimates of the age of basement rocks of the IBM forearc (Bloomer et al., 1995; Cosca et al., 1998) suggest that initial subsidence of the lithosphere along the IBM arc began at about 50 Ma. Depleted tholeiites and boninites erupted at the forearc at this stage (Hickey and Frey, 1982; Stern et al., 1991; Taylor et al., 1994). Subsidence of the lithosphere evolved into subduction at about 43 Ma (Eocene) when the motion of the Pacific plate changed abruptly from a northerly to a westerly direction (Richards and Lithgow-Bertellone., 1996).

The magmatic arc was close to its present position at the commencement of subduction (i.e., the distance from the magmatic arc to the trench has changed little since the formation of the magmatic arc). Single-arc volcanism continued until rifting started along the

Parece Vela Basin at 30 Ma (Oligocene). At 25 Ma, the Shikoku basin started spreading at the northernmost end of the IBM arc and propagated to the south (Kobayashi et al., 1995; Okino et al., 1999). The two rifting systems met at 20 Ma, and spreading stopped at about 15 Ma. The Kyushu-Palau Ridge (KPR) was formed at the western edge of both basins. At the same time (15 Ma) the development of a subduction zone at the Nankai Trough initiated the collision of the northernmost part of the Izu arc with the central Japanese landmass. Back arc rifting along the Mariana Trough started at 10 Ma, and the consequent seafloor spreading began at 3–4 Ma (Bibee et al., 1980; Yamazaki and Stern, 1997). Intra-arc rifting west of the present-day volcanic front of the Izu arc started at 2 Ma.

In the rear arc side of the Izu-Bonin arc, across-arc seamount chains are observed. Volcanism began along the across-arc seamount chains and at adjacent isolated seamounts at ca. 17 Ma, slightly before the Shikoku Basin ceased spreading, and continued until ca. 3 Ma (Ishizuka et al. 1998; 2003). Volcanism along these chains occurred sporadically along their total length, although volcanoes in the western part of the chains ceased to erupt earlier than those to the east. Volcanism

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along the seamount chains ceased at the initiation of rifting behind the volcanic front at ca. 2.8 Ma (Ishizuka et al. 2002).

Remarkable geochemical variations along the across-arc seamount chains have been observed (e.g., Hochstaedter et al., 2000). These across-arc variations imply that: 1) the proportion of slab-derived fluid component becomes smaller toward the backarc; 2) fluid is the dominant slab-derived component in the volcanic front, while sediment melt becomes more significant in the rear-arc seamounts; 3) heterogeneity exists in the mantle wedge prior to the addition of slab-derived components; 4) decompression melting may predominate toward the rear-arc whereas flux melting may predominate toward the volcanic front, even without backarc spreading; and 5) felsic magmatism inherits these differences. A more enriched source, in terms of the less-fluid-mobile incompatible elements, is expected to underlie the across-arc seamount chains. It is unknown whether these patterns occur uniformly across the arc or only along the across-arc seamount chains – that is, whether slab-derived components add continental-type crust only in certain parts of the arc. In order to examine formation mechanism of the across-arc seamount chains

and their geochemical variations, we carried out two active source seismic cruises along and across the across-arc sea mount chains in 2006.

2. Experiments

2.1 KR0601

Wide-angle seismic data were acquired along the Minami-Ioujima ridge, which is located at the southern end of Izu-Bonin arc, by using a seismic system on R/V Kairei (Fig. 1). The cruise was from 5th to 27th of January. We deployed Ocean bottom seismographs (OBS) from 7th to 13th. Air-gun shooting was done from 14th to 17th. OBSs were recovered from 17th to 22nd. One OBS (Site 12) was not recovered. Acoustic transponder seemed to work, but due to unknown reason this OBS did not released. After recovering the OBS, air-gun shooting test with a flip-flop mode was done. Although we planned to acquired multichannel reflection seismic data in adjacent profiles, those profile have not been made due to bad sea condition in the beginning of the cruise (Table 1 and 2). This cruise was the first R/V Kairei's wide-angle seismic experiment using more than 100 OBSs.

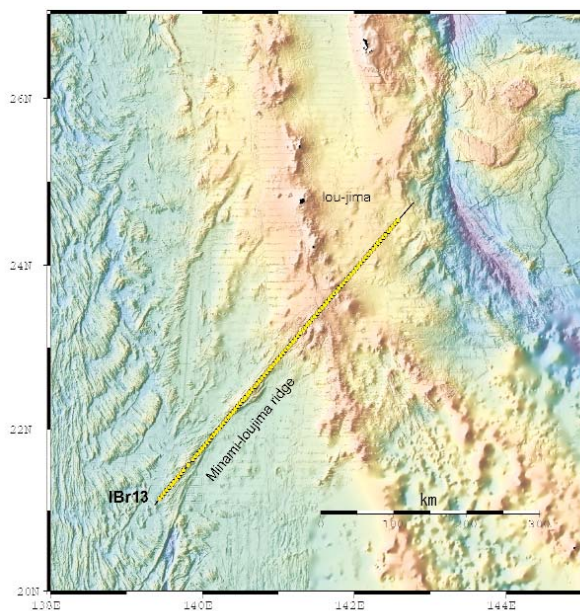


Figure 1: Profile map of KR0601

Table 1: Daily log of KR0601 cruise.

Date	Remarks
1/5	Departure from Jamstec
1/6	Transit
1/7	OBS Deployment (Site 100 – 91)
1/8	OBS Deployment (Site 90 – 71)
1/9	OBS Deployment (Site 70 – 48)
1/10	OBS Deployment (Site 47 – 30)
1/11	OBS Deployment (Site 29 – 19). Standby from 11:00
1/12	OBS Deployment (Site 18 – 10). Standby from 17:00
1/13	OBS Deployment (Site 9 – 1). Standby from 15:00
1/14	8:00. Deployment air-gun. 11:30. Start shooting IBr13
1/15	Shooting IBr13
1/16	Shooting IBr13
1/17	4:00 Stop shooting. end of IBr13. 7:00 OBS Retrieval (Site 100 – 92)
1/18	OBS Retrieval (Site 91 – 69)
1/19	OBS Retrieval (Site 68 – 48)
1/20	OBS Retrieval (Site 47 – 31)
1/21	OBS Retrieval (Site 30 – 19). Recovery drifting OBS (Site 14)
1/22	OBS Retrieval (Site 18 – 1). Site 2 was not recovered.
1/23	Flip-flop shooting test. 11:00 Retry Site 2, not recovered.
1/24	Retry Site 2, not recovered.
1/25	Transit
1/26	Transit
1/27	Arrival at Jamstec

Table 2: Profile positions of KR0601

Line ID	Date	Time (JST)	Position		SP
			Lat (N)	Lon (E)	
IBr13	2006/1/14	11:35:33	21°04'24.274"N	139°22'35.378"E	996
	2006/1/17	03:53:19	24°45'25.355"N	142°47'04.719"E	3682

2.2 KR0605

Wide-angle seismic data across the Nishi-shichito ridge between 27° N and 32° N and multichannel reflection data along two profiles from the Kinan seamount chain to the western edge of the Nishi-shichito ridge were acquired during this cruise (15th of May to 6th of June) (Fig. 2, Table 3 and 4). OBSs were deployed from 16th to 21st. Air-gun was shot for the OBSs to acquired the wide-angle seismic data from 21st to 25th. All OBSs were recovered from 26th to 31st. An acoustic transponder mounted on an OBS (Site22) did not respond, but it came up to sea surface without any reply. The multichannel reflection data were acquired from 1st to 4th of June. Sea condition was very stable during the entire period of this cruise. A new type of an acoustic transponder (System-Giken) is used in this cruise. It is attached to 27 OBSs. The new system, which is directory controlled by the ship-board acoustic system, worked fine except “inactive” command. The “inactive” command to stop replying seemed to be automatically reset sometime after this command was accepted.

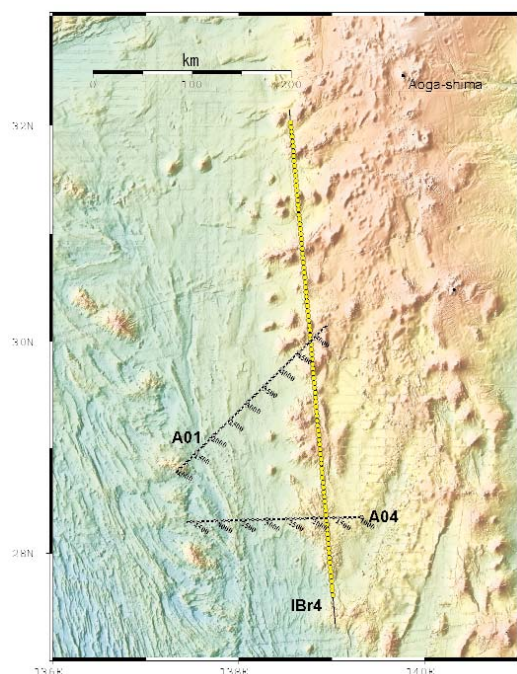


Figure 2: Profile map of KR0605

3. Seismic systems

3.1 Air-gun array

An air-gun array equipped on R/V Kairei was used for all profiles during the cruise. Shooting interval was 50 m, corresponding 20 – 30 s with accuracy of shot time of 1 ms for acquiring multichannel reflection data, while 200 m for acquiring wide-angle seismic data. Total volume of the air-gun array we used was 12,000 cubic inches, which consists of eight Bolt long life air-gun with 1,500 cubic inches each. The standard air pressure was 2000 psi (140 MPa). As shown in Fig. 3, two string of sub-array were deployed at the port and starboard sides of the vessel. Their width was expanded to 88.0 m by paravane system and the central position of the array was set 197.0 m behind the ship GPS antenna position.

3.2 Ocean bottom seismograph

For recording the wide-angle seismic data, we deployed 100 OBSs with 5 km interval during both cruise (Table 5 and 6), and recovered all of them except

Table 3: Daily log of KR0605

Date	Remarks
5/15	Departure from Yokosuka-Shinkou port
5/16	OBS Deployment (Site 1 – 10)
5/17	OBS Deployment (Site 11 – 30)
5/18	OBS Deployment (Site 31 – 48)
5/19	OBS Deployment (Site 49 – 70)
5/20	OBS Deployment (Site 71 – 90)
5/21	OBS Deployment (Site 91 – 100), 19:00 Start shooting IBr4
5/22	Shooting IBr4
5/23	Shooting IBr4
5/24	Shooting IBr4
5/25	3:00 Stop shooting, end of profile. 12:00 OBS Retrieval (Site 1 – 9)
5/26	OBS Retrieval (Site 10 – 23). Site22 did not respond
5/27	OBS Retrieval (Site 100 – 93)
5/28	OBS Retrieval (Site 92 – 74)
5/29	OBS Retrieval (Site 73 – 52)
5/30	OBS Retrieval (Site 51 – 34)
5/31	OBS Retrieval (Site 33 – 24). Retry Site22, recovered
6/1	7:00 Deployment air-gun. 15:10 Start shooting A01
6/2	17:00 Stop shooting, end of A01. 23:00 Start shooting A01
6/3	Air-gun shooting
6/4	6:30 Stop shooting, end of A01.
6/5	Transit
6/6	Arrival at Harumi port

Table 4: Profile positions of KR0605

Line ID	Date	Time (JST)	Position		SP
			Lat (N)	Lon (E)	
IBr13	2006/1/14	11:35:33	21°04'24.274"N	139°22'35.378"E	996
	2006/1/17	03:53:19	24°45'25.355"N	142°47'04.719"E	3682

for one OBS which was used in KR0601 (Site 22). The OBSs used (4.5 Hz, three-component gimbal-mounted geophones and hydrophones, continuous 16-bit digital recording with 100 Hz) were originally designed by Kanazawa and Shiobara (1994) and Shinohara et al. (1993). Continuous round motions were recorded on digital audiotape or a hard disk. Rechargeable lithium-ion batteries are used as an electric power supply. All parts including sensors with gimbal-leveling mechanism, batteries and a recorder are installed in 17-inch glass spheres made by Benthos, Inc, USA and Nautilus Marine service GmbH, Germany. An OBS equipped an acoustic transponder system which control release mechanism as well as positioning system. An OBS is released from anchors by electric corrosion of stainless plates when a release command was sent from the vessel. For positioning an OBS, super short base line (SSBL) acoustic positioning system is used. An accuracy of the SSBL system is about 3 % of water depth. After correcting a drift of an OBS clock by comparing the OBS clock and GPS, 70 s long data for each shots are formatted with SEG-Y.

3.3 Streamer cable

For acquiring the multichannel reflection seismic data during KR0605, we used a 204-channel hydrophone streamer cable made by Sercel Inc. (Fig. 4) The streamer cable is composed of 68 active sections, and each active section is 75m long and consists of three receiver groups (channels). The active modules including 24bit A/D converters are inserted every four active sections and collect seismic data from the four sections. The interval of each group is 25m. The lengths of total active section and lead-in cable are 5100m (75m × 68) and 110m, respectively. Hydrophone sensors (Benthos Reduce Diameter Array hydrophone) with sensitivity of 20V/Bar are used and the signals from 32 sensors in the same group (channel) are stacked before A/D conversion. The towing depth of streamer cable was controlled to be 15m below sea surface by the depth controller called Bird (DigiCOURSE System3).

3.4 Recoding and navigation system

The recording system of multichannel reflection data is the SYNTRAK960-24, which was made by Sercel Inc., and outputs seismic data onto 3590E tapes with SEG-D 8048 format. We set system delay to be 150ms. The sampling rate was 4ms and the record length was 15s. The Differential GPS (DGPS) was used for the positioning. We adopted StarFire system as a main positioning system and SkyFix as a backup. The accuracy is reported to be about 0.4 m in StarFire and 5 m in

SkyFix. As navigation software for seismic data acquisition, we used the SPECTRA made by Concept Inc. Positioning data collected from StarFire as well as SkyFix were sent to RTN μ (a network interface made by Concept Inc.) via a terminal server connected to LAN in the vessel. The RTN μ obtains time signals of DGPS (StarFire) from the original antenna. Then, the navigation data is sent to the PC Linux machine, on which the SPECTRA was installed and displayed. Shot times and Shot Point (SP) are set on the SPECTRA and then a trigger signal is sent to the recording system and the gun controller (GCS90), as follows.

First, a system-start-signal generated from the SPECTRA is sent to the recording system via the RTN μ , and soon after, the recording system send back a reply signal to the SPECTRA when the system is ready for recording. Second, the SPECTRA sends a trigger signal to the gun controller and the recording system sends a data-acquisition-start-signal to the streamer cable. The gun controller sends back an internal-time-break-signal to the RTN μ and recording system, and also sends trigger-signals to the eight airguns as shooting orders just after receiving the trigger signal from the SPECTRA. At the same timing with shot, the gun controller starts to gather both position data of the airgun sub-arrays from the airgun positioning system (RGPS) and first breaks of near-field shot records from monitor hydrophones nearby guns. Then seismic data are transmitted from the active modules to the recording system and position data of the streamer are sent from the depth controllers (so-called birds). After that, the position data of both the airgun and streamer cable are stored into the SPECTRA via the RTN μ and are also sent to the recording system. Finally, the seismic data are output to a tape drive and recorded on 3590E tapes. The recording system and gun controller are connected via RTN μ as shown by Figure 5.

4. Example of data

4.1 Wide-angle seismic data along the Minami-Ioujima ridge (KR0601)

All seismic section shown in Fig. 5 are plotted with 3 – 12 Hz band-pass filter, predictive deconvolution filter and 1s automatic gain control.

Site01: Quality of data is relatively good. First arrival can be traced up to 300 km offsets. This section is characterized by faster apparent velocity at near offset. The apparent velocity of ~ 8 km/s is observed from only 10 km offsets. This implies remarkably thin crust at the southwestern end of the ridge. Although clearer later reflections are hardly observed, a strong amplitude between 100 and 200 km offset may correspond a sub-moho reflection.

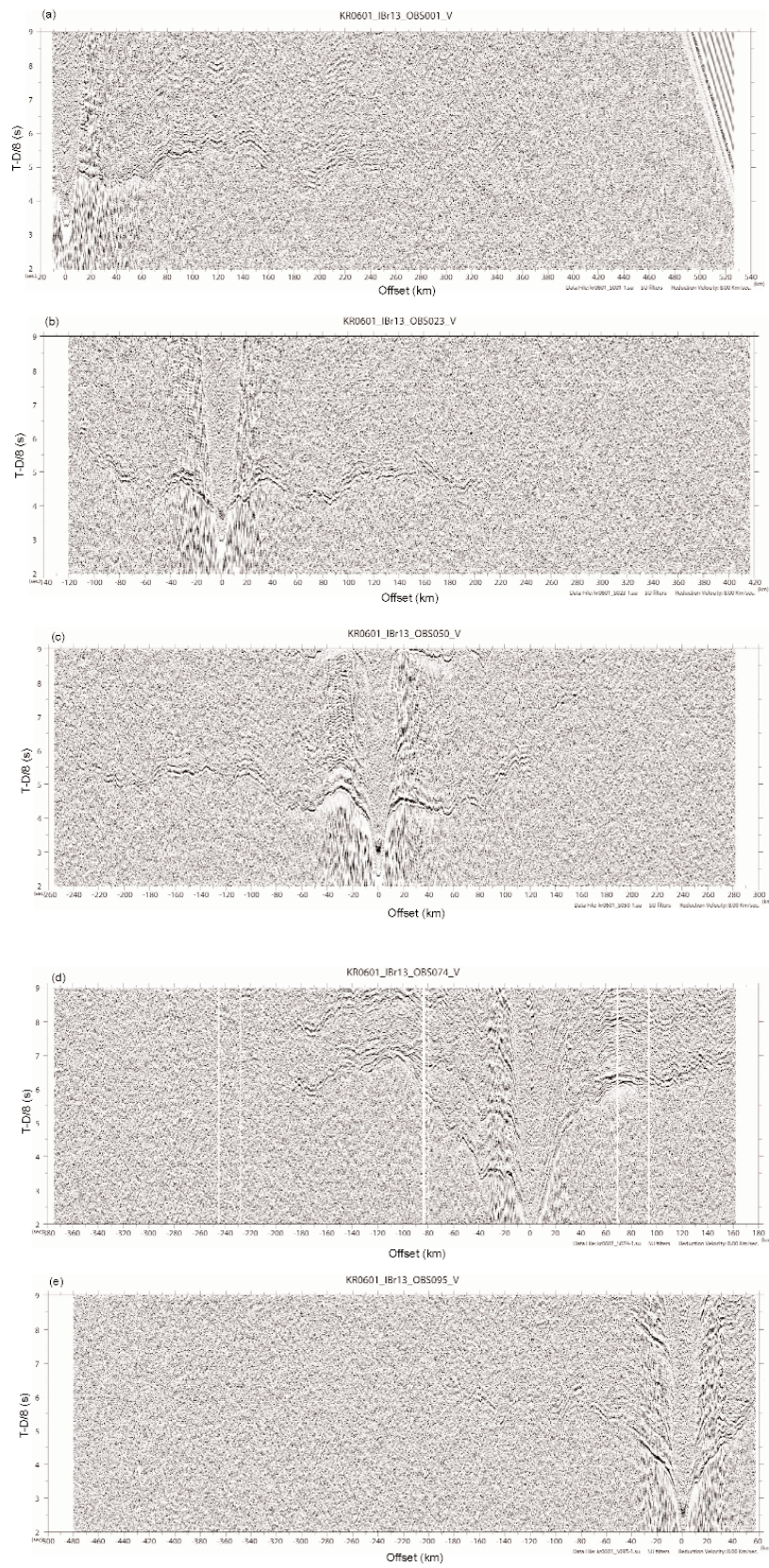


Figure 5: Example of wide-angle seismic data from KR0601. (a) Site01, (b) Site23, (c) Site50, (d) Site74, (e) Site90.

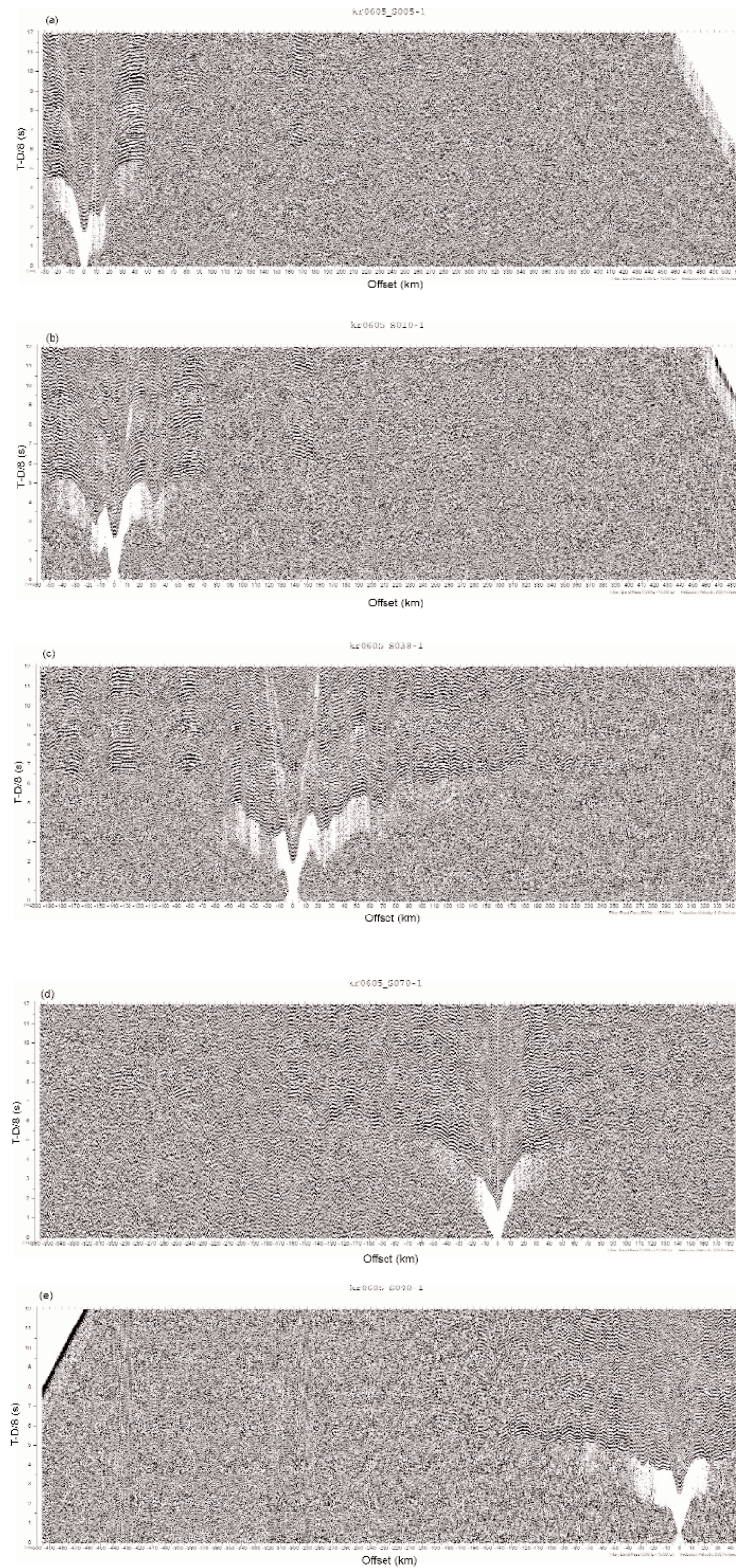


Figure 6: Example of wide-angle seismic data from KR0605. (a) Site05, (b) Site10, (c) Site38, (d) Site70, (e) Site98.

Site23: We can trace first arrivals more than 200 km offset from this record section. In contrast with the data observed at Site01, first arrivals showing apparent velocity less than 6 km/s are observed at shorter offset (~ 40 km offsets). Clear later reflection arrivals are recognized between 160 – 200 km offsets at 4.5 – 5.5 s. Those indicate an existence of a sub-moho reflector. Since the reflection from sub-moho reflector is also observed at Site01, this reflector may extend at the southern part of the ridge.

Site50: First arrivals can be traced more than 250 km offset at the southwestern part of the profile, while they can be traced up to only 160 km offset at the northeastern part. Characters of the first arrivals in the southwestern part are similar to Site 23. This indicates simple flat structure at least in the southwestern half of the profile. The apparent velocity of the first arrival from the northeastern part is distorted due to abrupt change of water depths. Stronger amplitude observed at 90 – 130 km offsets and 140 – 170 may indicate reflection phases.

Site74: This record section shows remarkably different character from the previous record sections. The most significant difference is longer observable range of slower apparent velocity phases (> 6 km/s). Those phases are observed up to 100 km offset at the southwestern part and 60 km offsets at the northeastern part. Since this OBS is located at near the top of the arc, those longer observable range indicate significantly thicker crust beneath the arc than the southern part of the ridge. Later arrivals observed at 70 – 15 km offsets (7 s) at the southwestern part and 60 – 120 km offsets (6.5 s) at the northeastern part are interpreted as reflections from the base of the crust beneath the arc. We can estimate exact thickness of the arc crust by further processing.

Site95: Observable ranges of slower apparent velocity less than 6 km/s is shorter than that of Site74. This implies thinning of the crust toward forearc. A weaker amplitude phase observed more than 80 km offsets may correspond mantle reflection phase. Although later reflection phases are not very obvious, larger amplitude phases at 40 – 70 km and 80 – 120 km offset may represent reflection phases.

4.2 Wide-angle seismic data across the Nishishichito ridge (KR0605)

Seismic sections shown here are only process with 5 – 15 Hz band-pass filter and automatic gain control of 2 s as preliminary processing (Fig. 6).

Site05: Since the profile is designed for across the ridges, apparent velocities of first arrivals are strongly affected by variation of water depth along the profile (e.g., 5 – 20 km offsets and 50 – 70 km offsets). First

arrivals observed farther than 30 km, however, roughly show apparent velocity of 8 km/s which may be refraction arrival from mantle. Stronger amplitude at 80 – 110 km and 160 – 200 km offset may be reflections phases.

Site10: This record section shows good quality of arrivals. We can trace the signal up to 250 km offset. The slower apparent velocity phase less than 7 km/s observed till 70 km offset is interpreted as crustal refraction phases. Later arrivals at 50 – 90 km and 14 – 200 km offsets may be reflections from the base of crust and sub-moho reflector, respectively.

Site38: This record section also shows good quality data. First arrival can be traced more than 250 km offset. Slower apparent velocity phases less than 7 km/s are observed 120 km at the south and 80 km at the north. They shows longer offsets than those observed at Site10. This implies that crust becomes thicker toward the middle of the profile from the north. Strong amplitude later arrivals observed from 80 km offset at both side are interpreted as reflections from the base of crust. Another later arrival observed at 120 – 150 km offset may correspond to reflection from a sub-moho reflector at the northern part of the profile.

Site70: Observable range of slower apparent velocity phase (< 7 km/s) becomes shorter in this section than those of site 38. This indicates crustal thinning toward the south from the middle of the profile. Later arrivals observed at 40 – 60 km offset may be reflection from the base of crust.

Site98: First arrivals interpreted as crustal refraction phases are observed up to 130 km offset. Apparent velocity of those phase seem to be firster than the previous sections, but this may be cause of variation of water depth. Due to reverberation of the first arrival, onset of later reflection arrivals are not clearly recognized, however, wave-train observed between 50 and 120 km offset seems to show reflections from the base of crust.

4.3 Multichannel reflection seismic data

Raw MCS reflection data are processed on board for the purpose of quality control and preliminary interpretation of tectonic structures in the study areas. The on-board data processing was conducted preserving relative amplitudes under the conventional processing scheme, as shown in Figure 7, which contains noisy-trace editing, 5-100 Hz band-pass filtering, deconvolution with a 28-ms-length predictive distance and a 250-ms-length operator, amplitude compensation by T^2 (T is two way traveltime), velocity analysis, multiple suppression by radon transform, muting, CDP stacking and post-stack time migration.

Figure 8 show results of on-board processing A01

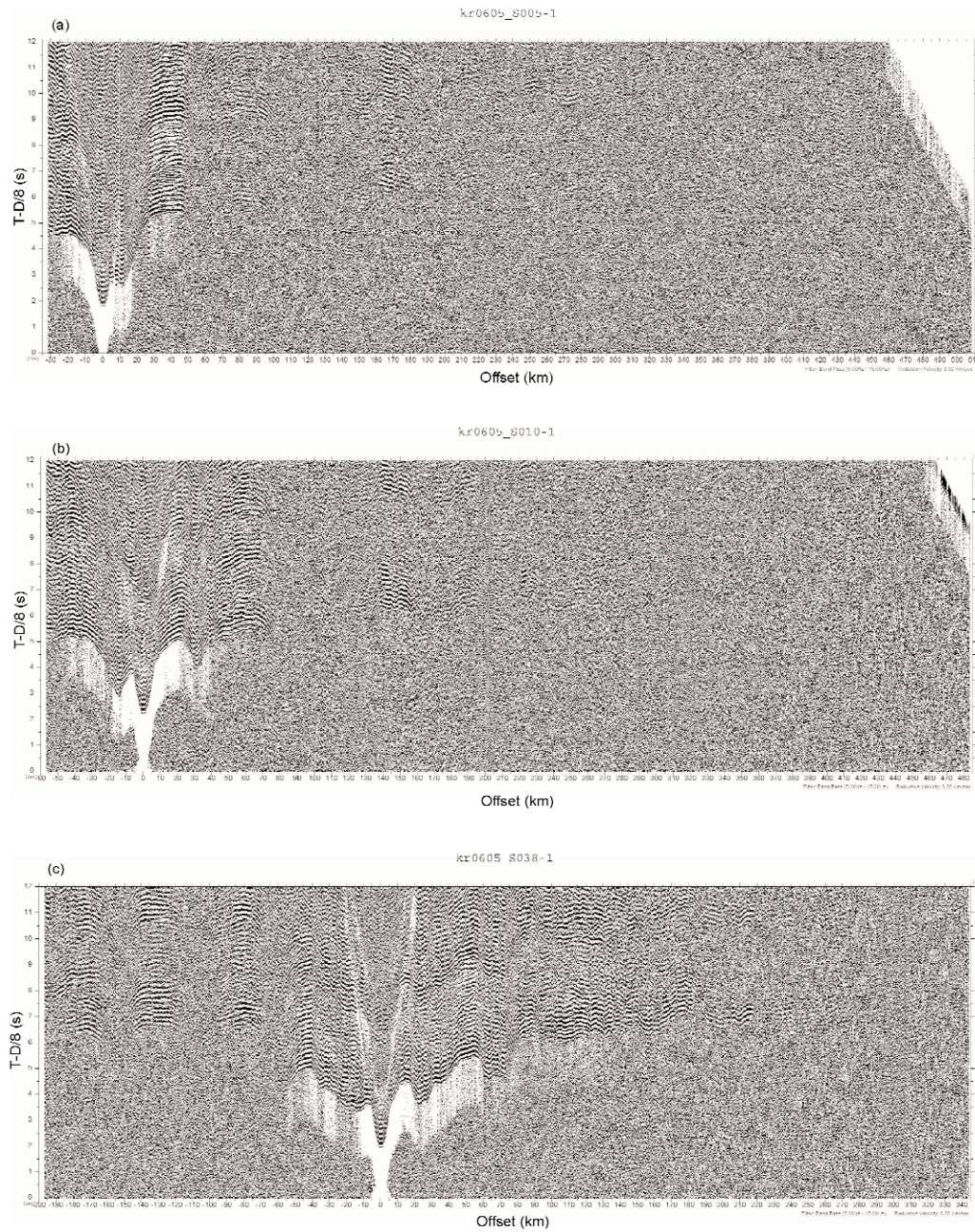


Figure 7: Flow chart of on-board processing multichannel data

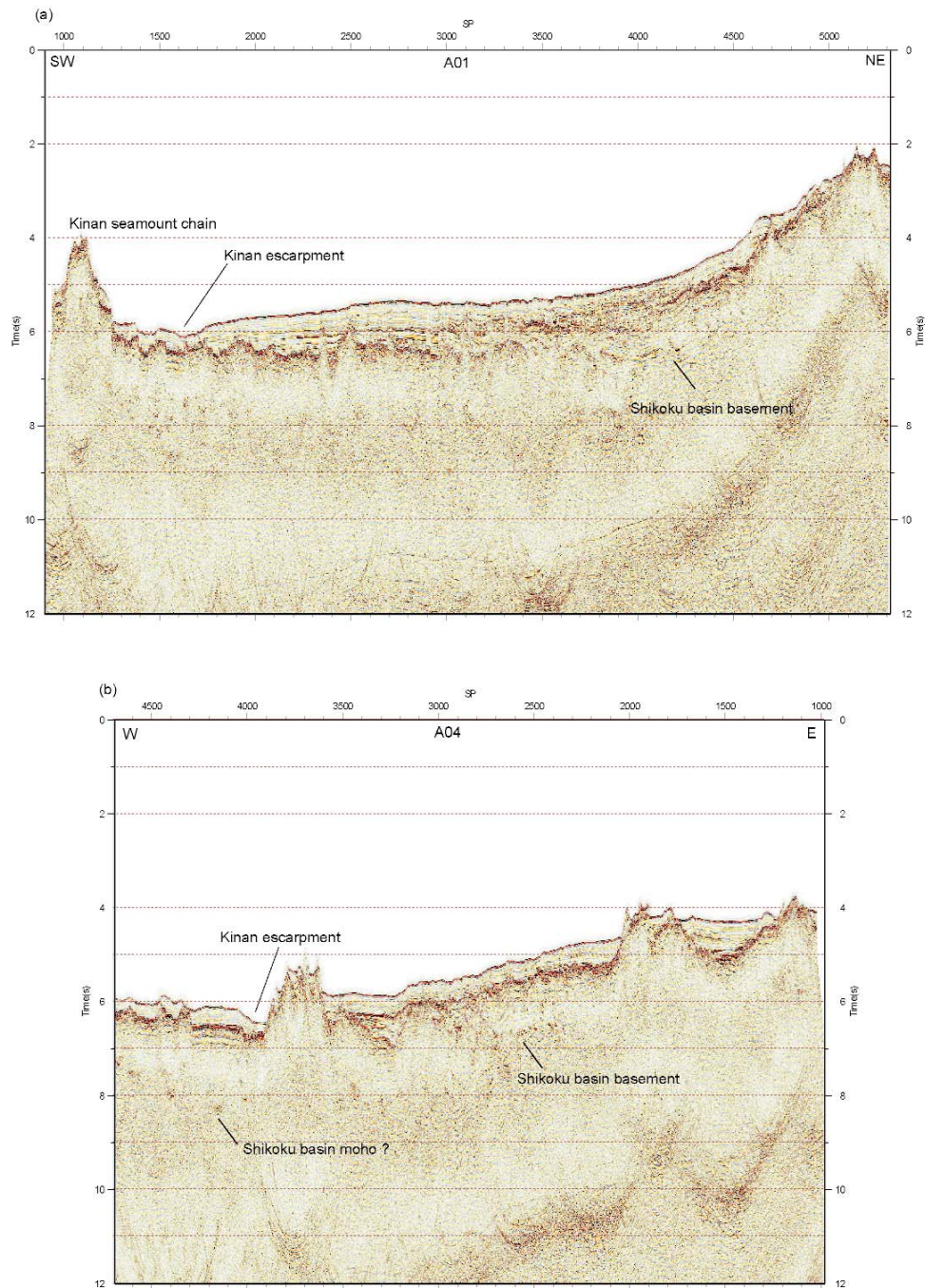


Figure 8: Results of on-board processing multichannel data. (a) Time migration of A01, (b) Time migration of A04.

and A04. Both profiles are located from the Kinan seamount chains to the Nishi-shitito ridges through the Kinan escarpment, A01 was take more parallel to the strike of across-arc seamount chains. The most characteristic structure in both sections is that the basement of the Shikoku basin is imaged beneath the basement extending from the across-arc seamount chains. We can recognize this character clearer along A01. In Fig.8, the Shikoku basin basement is traced at SP4300. The basement from the across-arc seamount chains cover the Shikoku basin basement till SP2500. This character indicates that the across-arc seamount chains were formed on the Shikoku basin crust, and a boundary between the Shikoku basin and the Oligocene arc may be situated around eastern end of the across-arc seamount chains. It notes that the base of the Shikoku basin crust is sporadically observed along both profiles. Although the Kinan escarpment has been suggested as a sort of structure boundary, any distinct structural variation beneath the basement can not be observed in A01.

5. Summary

We have conducted two active source seismic cruises across and along the across-arc seamount chains in the Izu-Bonin arc by R/V Kairei in 2006, KR0601 and KR0605. They are the first cruises in which more than 100 OBS were handled by R/V Kairei. We have successfully acquired wide-angle seismic data during both cruise and multichannel reflection data during KR0605. The observed wide-angle seismic data acquired along the Minami-Ioujima ridge (KR0601) indicate flat and thin crust beneath the ridge. The wide-angle seismic data across the across-arc seamount chain in the Izu arc show the north – south structural variation (i.e. thicker crust in the north and thinner crust in the south). We found that the basement of the Shikoku basin is continued under the across-arc seamount chain. This indicates that a boundary between the Shikoku basin crust and the Oligocene arc crust is situated around eastern end of the across-arc seamount chains.

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