

1. Introduction

Global warming resulting from the increase in greenhouse gases, such as carbon dioxide, is of great concern to the world community. In the last few decades, the carbon cycle in the ocean has been studied to clarify the balance of carbon dioxide between the atmosphere and the ocean. One important issue is the quantification of the role played by the biological pump: how much atmospheric CO_2 is assimilated in the sunlit layer (euphotic zone) and how much carbon is exported to the deep ocean?

The northern North Pacific Ocean, especially its western part, experiences intense winter cooling and receives large supplies of nutrients through upwelling, resulting in high productivity in spring and summer. It is well documented that spring blooms, consisting mainly of diatoms, occur only in the western part of the subarctic Pacific (Saito *et al.*, 2002; Yamaguchi *et al.*, 2002). It has been suspected that diatoms in this area play a key role in transporting POC to the deep ocean, because of their relatively large size and resultant high settling velocity (*e.g.*, Tsunogai and Noriki, 1991; Kemp *et al.*, 2000; Smetacek, 2000; Honda *et al.*, 2002). The large decrease of nutrients and surface $p\text{CO}_2$ from spring to autumn also support the high activity of the biological pump in the northwestern North Pacific (Louanchi and Najjar, 2000; Wong *et al.*, 2002; Takahashi *et al.*, 2002). Kawakami and Honda (2007) reported that seasonal variability in POC fluxes estimated from ^{234}Th was large ($54\text{--}179 \text{ gC m}^{-2} \text{ d}^{-1}$) and annual POC fluxes were estimated to be $31 \text{ gC m}^{-2} \text{ y}^{-1}$ in this region.

^{230}Th and ^{231}Pa are useful tracers of sinking particle flux in the ocean, and the trapping efficiency of the sediment trap is estimated from these radionuclides (*e.g.*, Bacon *et al.*, 1985; Yu *et al.*, 2001).

We have made time-series sediment trap experiments in the subarctic western North Pacific since 2001 (Honda *et al.*, 2006 and 2009). We determined ^{230}Th and ^{231}Pa in the sinking particles collected by the sediment traps in 2005–2008 (*e.g.*, Shigemitsu *et al.*, 2010).

In this dataset, we present the fluxes of ^{230}Th and ^{231}Pa collected by the sediment traps in the western North Pacific. These data will help further understanding of particle dynamics in the ocean.

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3. Sampling and sample analysis

Time-series sediment traps (McLane Mark7G) with 13 or 21 collecting cups were deployed and recovered at 300, 540, 1000, and 4810 m depths at station K2 (47°N, 160°E; bottom 5200 m) in February–March 2005 (MR05-01, doi: 10.17596/0001792), September–October 2005 (MR05-04, doi: 10.17596/0001793), May–June 2006 (MR06-03_leg1, doi: 10.17596/0001796), September–October 2007 (MR07-05, doi: 10.17596/0001803), and October–November 2008 (MR08-05, doi: 10.17596/0001809). The station K2 is in the subarctic part of the western North Pacific Ocean. Collected samples were preserved with seawater based buffered 10% formalin. The pretreatment procedures of samples (storage, sieving, splitting, and filtration) followed the methods described by Honjo *et al.* (1995). A part of sinking particle samples was digested with a mixture of concentrated HNO₃, HClO₄, and HF (1:1:1, v/v) in the presence of the ²²⁸Th and ²³³Pa yield tracers (²²⁸Th: 5–10 dpm, ²³³Pa: 400–1000 dpm). Radiochemical separation and purification of these nuclides were achieved by procedures similar to those of Anderson and Fleer (1982) using the anion exchange technique. The treated Th and Pa were electroplated onto stainless-steel planchets (Narita *et al.*, 2003). The planchets for Pa were covered with polyethylene sheets (50 μm thick) and first beta-counted using a low-background (0.1–0.3 cpm) anticoincidence gas-flow beta detector (LBC-470, Aloka Co. Ltd.) to determine their ²³³Pa (yield tracer) activity, and then the planchets for Th or Pa were alpha-counted with silicon surface barrier detectors (Octéte, Seiko EG&G Co. Ltd.) to determine their ²²⁸Th (yield tracer), ²³⁰Th, ²³²Th, and ²³¹Pa activities. Excess ²³⁰Th and ²³¹Pa were estimated from following equations (Yu *et al.*, 2001):

$$\text{excess } ^{230}\text{Th} = ^{230}\text{Th}_T - 0.8 \cdot ^{232}\text{Th}_T$$

$$\text{excess } ^{231}\text{Pa} = ^{231}\text{Pa}_T - 0.046 \cdot 0.8 \cdot ^{232}\text{Th}_T$$

where ²³⁰Th_T, ²³⁰Th_T, and ²³¹Pa_T are the total activities measured. The precision of excess ²³⁰Th and ²³¹Pa was one sigma of the counting error.

3. Dataset

Data obtained from sediment trap experiments were electrically compiled as an excel file “ST_230Th_231Pa.xls”, and were total mass flux, flux of excess ^{230}Th (ex230Th), and excess ^{231}Pa (ex210Pa). The data of total mass flux were excerpted from “JAMSTEC OceanSITES Sediment Trap data (netCDF format)” (<https://ebcrpa.jamstec.go.jp/k2oceansites/>). The errors of ^{230}Th and ^{231}Pa were estimated from counting error.

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